

Towards a holistic understanding of complex magnetic fabrics

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Magnetic fabrics have initially been attributed to rock texture by Graham (1954) and Balsley and Buddington (1960), who found that the maximum susceptibility was parallel to the lineation and the minimum susceptibility normal to the foliation in a series of granitic rocks. Anisotropy of magnetic susceptibility (AMS) or remanence (e.g. anhysteretic, AARM, isothermal, AIRM, or thermal, ATRM) have since been used to infer lava flow directions and emplacement models in igneous rocks, deformation and tectonic transport directions in metamorphic rocks, and water-current or wind directions in sediments (cf. Borradaile and Henry, 1997; Borradaile and Jackson 2010; Hrouda 1982; Martín-Hernández et al., 2004; Owens and Bamford, 1976; Tarling and Hrouda, 1993), and the degree of anisotropy has been linked to the strength of deformation (Cogné and Perroud, 1988; Kligfield et al., 1977, 1981; Kneen, 1976; Parés and van der Pluijm, 2004; Weil and Yonkee, 2009). More recent studies have shown that anisotropy-deformation relationships need to be established for every specific rock



Beautiful folding in Oppdal, Norway. Photo: A. R. Biedermann

type, taking into account the mineralogy and the relative proportions of ferromagnetic minerals with respect to the paramagnetic and diamagnetic constituents of the rocks, for example (Borradaile and Henry, 1997; Housen and van der Pluijm, 1990; Housen et al., 1993). Furthermore, grain-size can influence magnetic anisotropy; for example single-domain (SD) magnetite carries “inverse AMS fabrics”, with minimum susceptibility parallel to the flow direction or lineation, opposite to the common interpretation (e.g. Gialanella et al., 1994; Rochette et al., 1999; Hrouda and Jezek, 2017). Still, magnetic fabrics are a powerful indicator of a rock’s texture, particularly the preferred alignment of grains, which is modified during geodynamic processes by grain re-alignment, grain-deformation, or recrystallization.

Rock texture is a complex property, involving the different mineral constituents and their crystallographic preferred orientation (CPO), shape preferred orientation (SPO), and spatial distribution in the rock. Whereas textures can be described directly using a variety of techniques like electron backscatter diffraction (EBSD; Prior et al., 1999), or image analysis (Heilbronner and Barrett, 2014), it is often easier, and more representative for heterogeneous materials, to assess preferred mineral orientation by measuring the anisotropy of physical properties, like seismic or magnetic anisotropy (Engler and Randle, 2009; Ullemeyer et al., 2000). Magnetic fabrics are described by a magnitude ellipsoid whose properties

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pg. 8.*

Visiting Fellow Report

A behavioral archeomagnetic investigation of burnt rock middens in central Texas.

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I visited the IRM to conduct a magnetic study of hot rock cooking technology on the Edwards plateau, which initiated around 8,000 BP. Use of hot rock cooking technology is extremely visible in the study area, often producing very large (>50m in diameter) mounds or 'middens' of used burnt rock. Goal of the project was to provide quantitative observations to identify specific behaviors associated with the formation and use of these structures.

The first scientific investigations of burnt rock middens in central Texas were conducted by J.E. Pearce from 1903 to 1938, who concluded that middens were comprised of cooking debris from hearth structures. In 1942 Kelly and Campbell instead suggested that middens are made up of remnants of stone-lined basins that were used for cooking. In 1969 Sorrow proposed the interpretation that middens were dump areas where broken hearth stones were discarded, but it was not until 1974 that Prewitt (n.d) argued for middens representing the remains of covered stone-lined earth ovens. The data I gathered at the Institute for Rock Magnetism (IRM) at the University of Minnesota will be used to further test these hypotheses at the Gault site.

During my time at the IRM I carried out thermal demagnetization of oriented cores collected from middens of the Gault site to investigate whether any post-heating movement of the individual limestone blocks had occurred (Fig. 1). I also carried out low temperature remanence measurements (Fig. 2), magnetic hysteresis (Fig. 3) and magnetic susceptibility as a function of temperature to characterize the mineralogies responsible for the magnetic remanence.

The information collected will be used to reconstruct the original morphology of a stone heating structure from the scattered debris that make up middens (i.e. simple hearths, stone-lined basins, or covered earth ovens), which directly addresses the aforementioned hypothesis.

One significant example about the formation of burnt rock middens comes from the Wilson Leonard site in Central Texas. Archaeologists were able to use archeomagnetic data from culturally heated limestone rocks to determine that middens at the Wilson Leonard site are likely aggrading living surfaces rather than large discard piles. Another significant discovery was that many of the stone-lined middens which were assumed to be open roasting or grilling pits, were actually earth ovens that were closed with a stone lid, as asserted by Prewitt. This

was determined by examining rocks directly adjacent to stone-lined pits. At many sites the majority of these rocks were moved away from the pit at a common low temperature, and are interpreted to be the lid of the ovens.

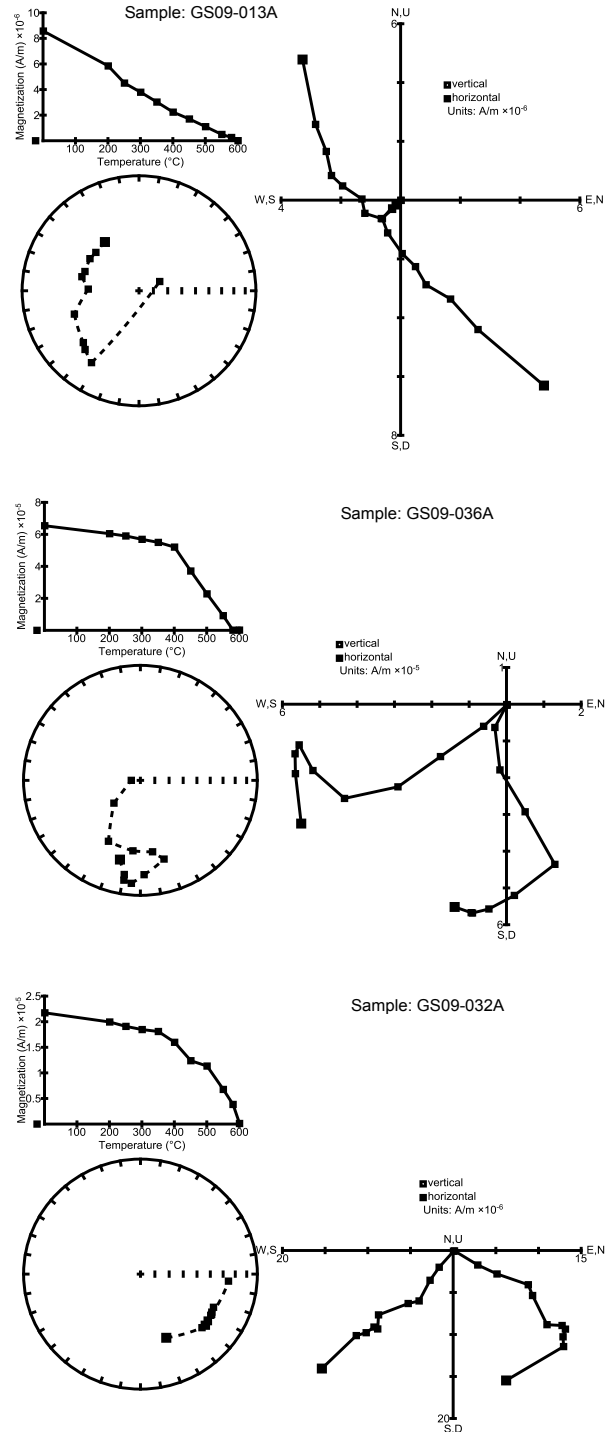


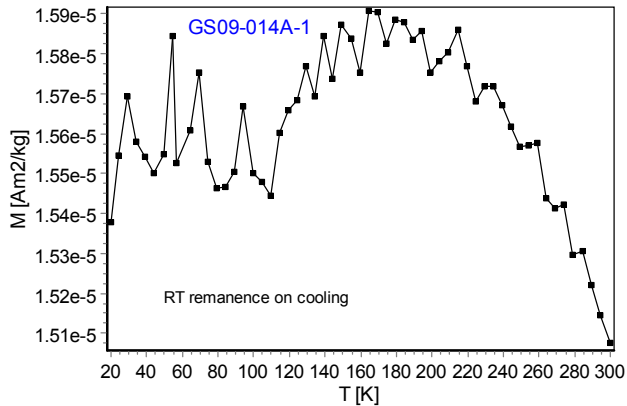
Figure 1. Thermal demagnetization of the NRM indicating three different demagnetization behaviors: a quasi constant decay of remanence between room and 600°C indicating a range of magnetite grain-sizes as dominant magnetic carrier, and two distinct magnetization components; a "narrower" range of magnetite unblocking temperatures between 400° and 580°C, and multiple superimposed magnetization components; and two dominant unblocking temperatures ~400° and 600°C, and at least two associated components of magnetization.

Current Articles

A list of current research articles dealing with various topics in the physics and chemistry of magnetism is a regular feature of the IRM Quarterly. Articles published in familiar geology and geophysics journals are included; special emphasis is given to current articles from physics, chemistry, and materials-science journals. Most are taken from ISI Web of Knowledge, after which they are subjected to Procrustean culling for this newsletter. An extensive reference list of articles (primarily about rock magnetism, the physics and chemistry of magnetism, and some paleomagnetism) is continually updated at the IRM. This list, with more than 10,000 references, is available free of charge. Your contributions both to the list and to the Current Articles section of the IRM Quarterly are always welcome.

Aeromagnetism, Magnetic Anomalies, and Surveying

- Anudu, G. K., R. A. Stephenson, D. I. M. Macdonald, and G. N. Oakey (2016), Geological features of the northeastern Canadian Arctic margin revealed from analysis of potential field data, *Tectonophysics*, 691, 48-64.
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- Xiao, F., and Z. H. Wang (2017), Geological interpretation of Bouguer gravity and aeromagnetic data from the Gobi-desert covered area, Eastern Tianshan, China: Implications for porphyry Cu-Mo poly-



2. Low-temperature demagnetization of a 300K saturation isothermal remanent magnetization, showing a distinct magnetite Verwey transition over an overall weak magnetization. Goethite is also apparent from measurements of other specimens, not shown.

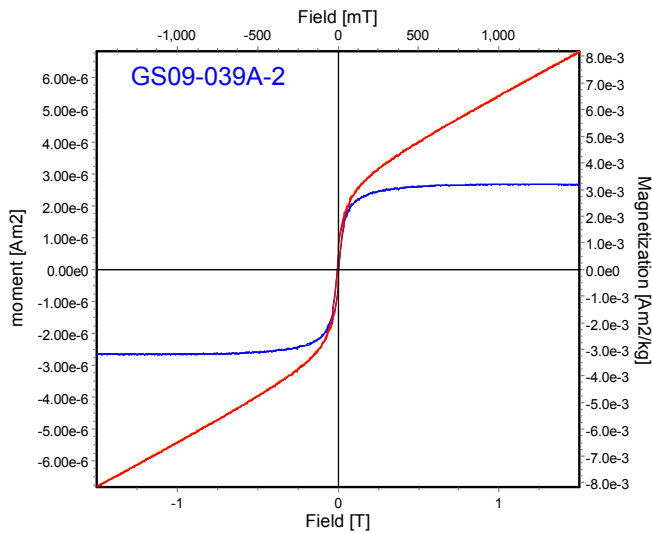


Fig. 3. Hysteresis loop indicating multi-domain magnetite dominant signal, and little wasp-waistedness indicating a mixture of coercivities present.

My magnetic findings from the Gault site will complement the observations of the Wilson Leonard site, further proving this new understanding of behaviors associated with the use of stone cooking structures, and the formation of burnt rock middens. Furthermore the identification of specific actions taken by individuals at a particular moments in time, will strengthen the case for behavioral archeology and place the role of the individual in the broader pictures painted by archeologists. Behavioral archeomagnetism in central Texas offers the rare chance to identify specific behaviors taken by individuals at a specific moment in time. Like preserved human footprints, the thermal history of a culturally heated rock is a trace fossil of human behavior.

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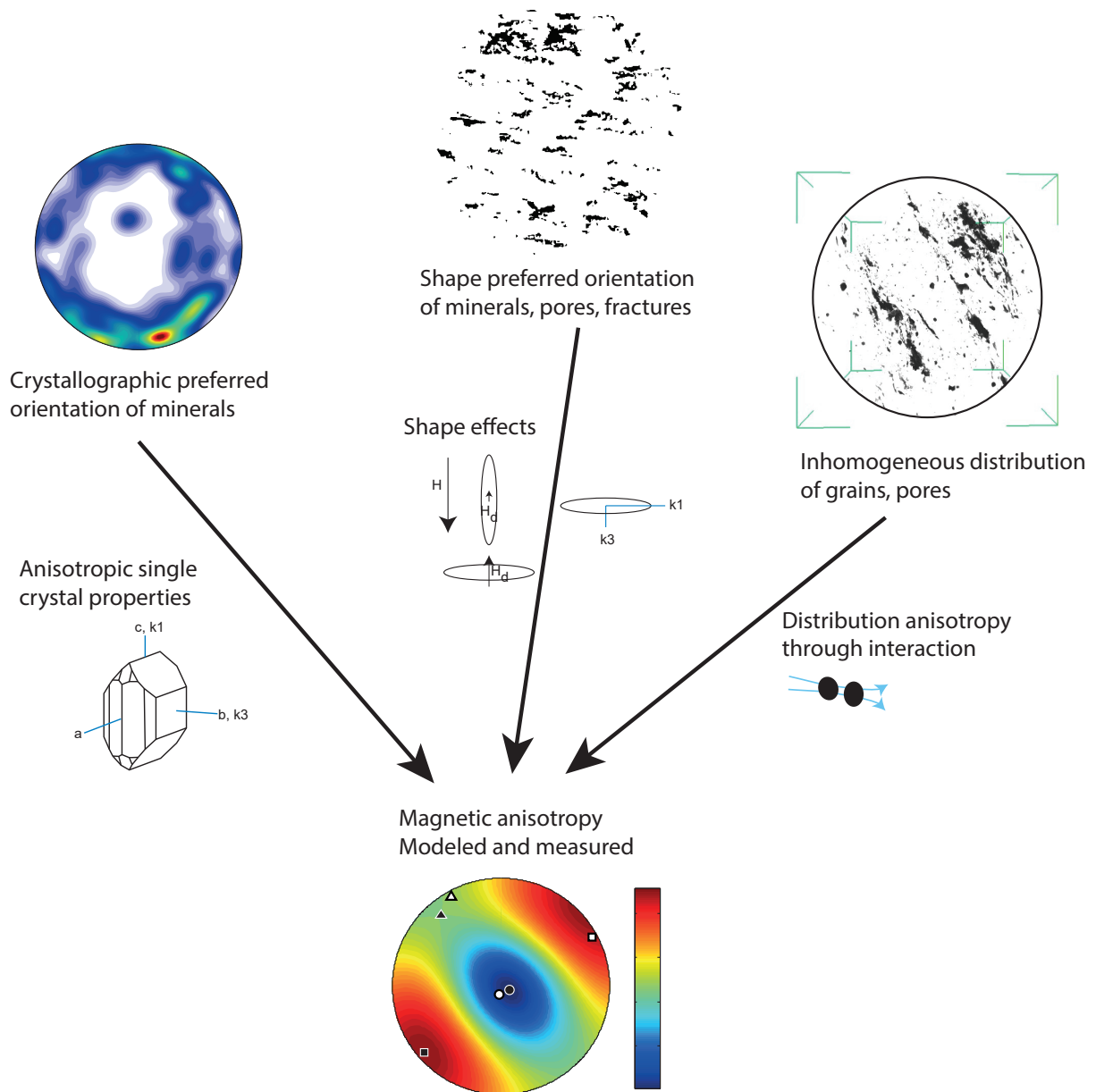


Fig. 1. Conceptual sketch of petrofabric properties influencing magnetic anisotropy in rocks. Please refer to online version for color.

relate to the more complicated rock texture.

Because the magnetic anisotropy ellipsoid is a simplified representation of the rock texture, challenges arise when interpreting the resulting fabrics, as in the case of inverse or complex fabrics. However, with a better understanding of factors controlling magnetic anisotropy, and improved experimental techniques that allow for isolating contributions of specific mineral groups, magnetic fabrics are now an even more reliable proxy for rock texture. Magnetic anisotropy is a direct consequence of (1) magnetocrystalline anisotropy in combination with CPO, (2) shape anisotropy together with SPO, and (3) distribution anisotropy and non-uniform spatial distribution of grains (Figure 1; Cañon-Tapia, 1996; Grégoire et al., 1995; Hargraves et al., 1991; Mainprice et al., 1994). Thus it is essential to understand (1) which mineral carries the anisotropy, (2) whether it possesses a magnetocrystalline, shape or distribution anisotropy, and (3) how the CPO, SPO and spatial distribution re-

late to geodynamic processes, before any interpretation of magnetic fabric data. Minerals carrying the magnetic anisotropy can be identified by (1) experimental separation techniques, e.g. measurements in various temperatures and fields, and anisotropy of ARMs and IRMs, to isolate the anisotropy carried by the paramagnetic and different types and grainsizes of ferromagnetic minerals (e.g. Martín-Hernández and Ferré, 2007; and references therein), and (2) comparison between the orientation and degree of the AMS ellipsoid to the properties of either CPO or SPO of a specific mineral. The latter is especially suited for samples whose magnetic fabric is dominated by one mineral. For example, AMS in phyllosilicate-bearing rocks has been found to directly reflect the CPO of the phyllosilicates (Chadima et al., 2004; Hirt et al., 1995; Lüneburg et al., 1999; Richter et al., 1993; Siegesmund et al., 1995). Ferré et al. (2005) investigated fabrics in a dunite and compared high-field AMS to olivine CPO. Other studies show AMS correlating with the SPO

of magnetite grains (Archanjo et al., 1995, 2002; Grégoire 1995, 1998; Launeau and Cruden, 1998). If several minerals contribute to the magnetic fabric, detailed and systematic characterization of single crystal susceptibility tensors as well as physical models for magnetocrystalline, shape and distribution anisotropy are necessary for a reliable interpretation.

Modeling magnetic anisotropy and the contribution of individual minerals

A relatively new approach was adapted from a method used to compute seismic anisotropy, and focuses on modeling contributions of individual rock-forming minerals to the magnetocrystalline anisotropy (Mainprice and Humbert, 1994). These models are based on CPO data and single crystal susceptibility tensors, and can be used to predict (1) the contribution of specific minerals to the whole-rock anisotropy, or (2) the sum of contributions for a group of minerals or all minerals in the rock that possess magnetocrystalline anisotropy. The simulated components of anisotropy can then be compared to the measured AMS, or any experimentally isolated component of the AMS measured on the same rock. In any case, well-defined susceptibility tensors of single crystals for each mineral in the rock are necessary for reliable models. Whereas the first studies on single crystal magnetic anisotropy were published over a century ago (e.g. Faraday, 1846; Finke, 1909; Tyndall, 1851), recent works have systematically characterized magnetocrystalline anisotropy as a function of chemical composition for many common rock-forming minerals from both the carbonate and silicate groups (e.g. Biedermann et al., 2014a, 2015a; Schmidt et al., 2006, 2007). Schmidt et al. (2009) compared the paramagnetic and diamagnetic components of the AMS in synthetic quartz-muscovite aggregates to those modeled based on muscovite and quartz CPO, respectively. Biedermann et al. (2015b) modelled the paramagnetic component of AMS in amphibolites, peridotite and pyroxenite, nicely matching the isolated paramagnetic component of AMS (Figure 2). Additionally, modeling allows to investigate the interplay between fabrics of different paramagnetic minerals which cannot yet be separated experimentally. Furthermore, using synthetic fabrics, it is possible to simulate under which circumstances the anisotropies of two minerals interfere positively, creating a stronger whole-rock magnetic fabric, or negatively, leading to an overall weaker fabric (Figure 3). Such modeling is clearly beneficial when interpreting complex fabrics.

Summary

Experimental separation techniques and understanding which minerals carry the anisotropy have significantly increased the reliability of the interpretation of magnetic fabrics. Together with new single crystal data, modelling of the anisotropy contribution carried by different minerals in a rock allows to take magnetic fabric interpretation one step further, particularly in the case of complex fabrics. A more detailed understanding of the interplay of different minerals to form the whole-rock

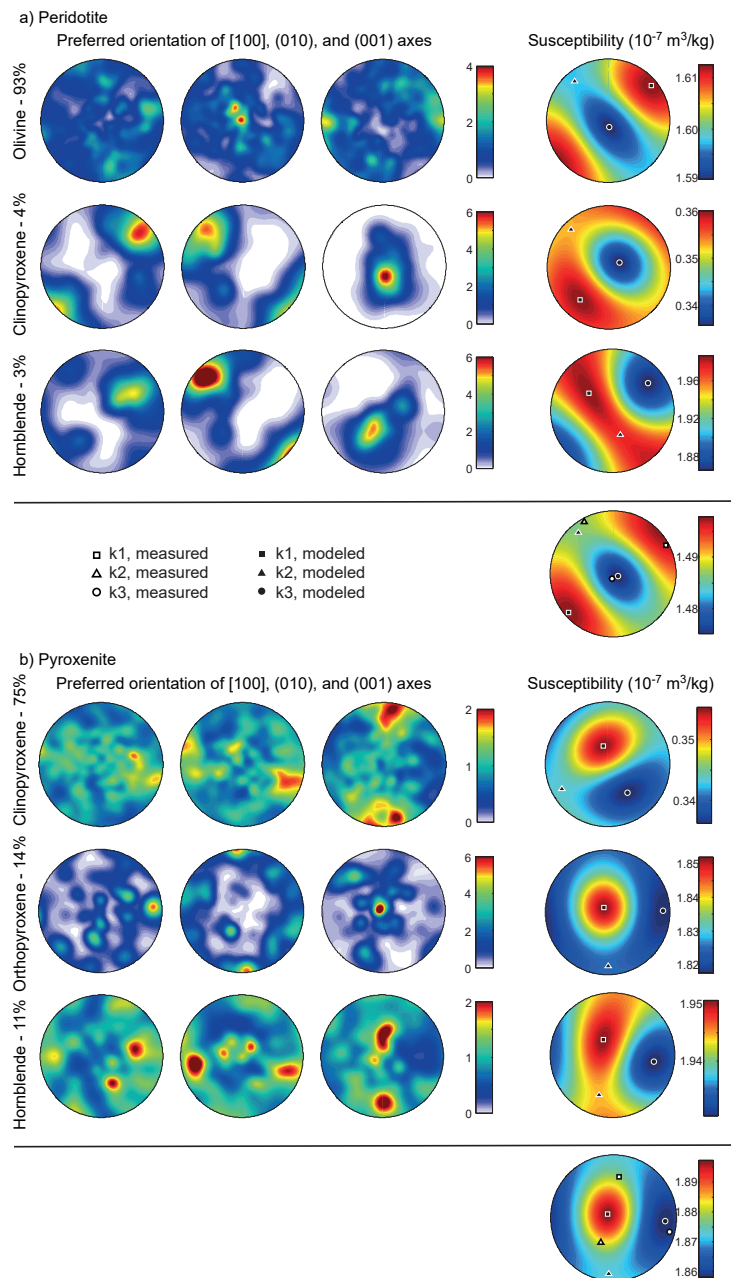


Fig. 2.(a) Peridotite with negative interference between olivine and hornblende AMS, and (b) pyroxenite with positive interference between orthopyroxene and hornblende AMS. Modified after Biedermann et al., 2015b.

magnetic anisotropy will provide a deeper insight into different AMS components, and eventually lead to improved interpretation of magnetic fabrics in rocks.

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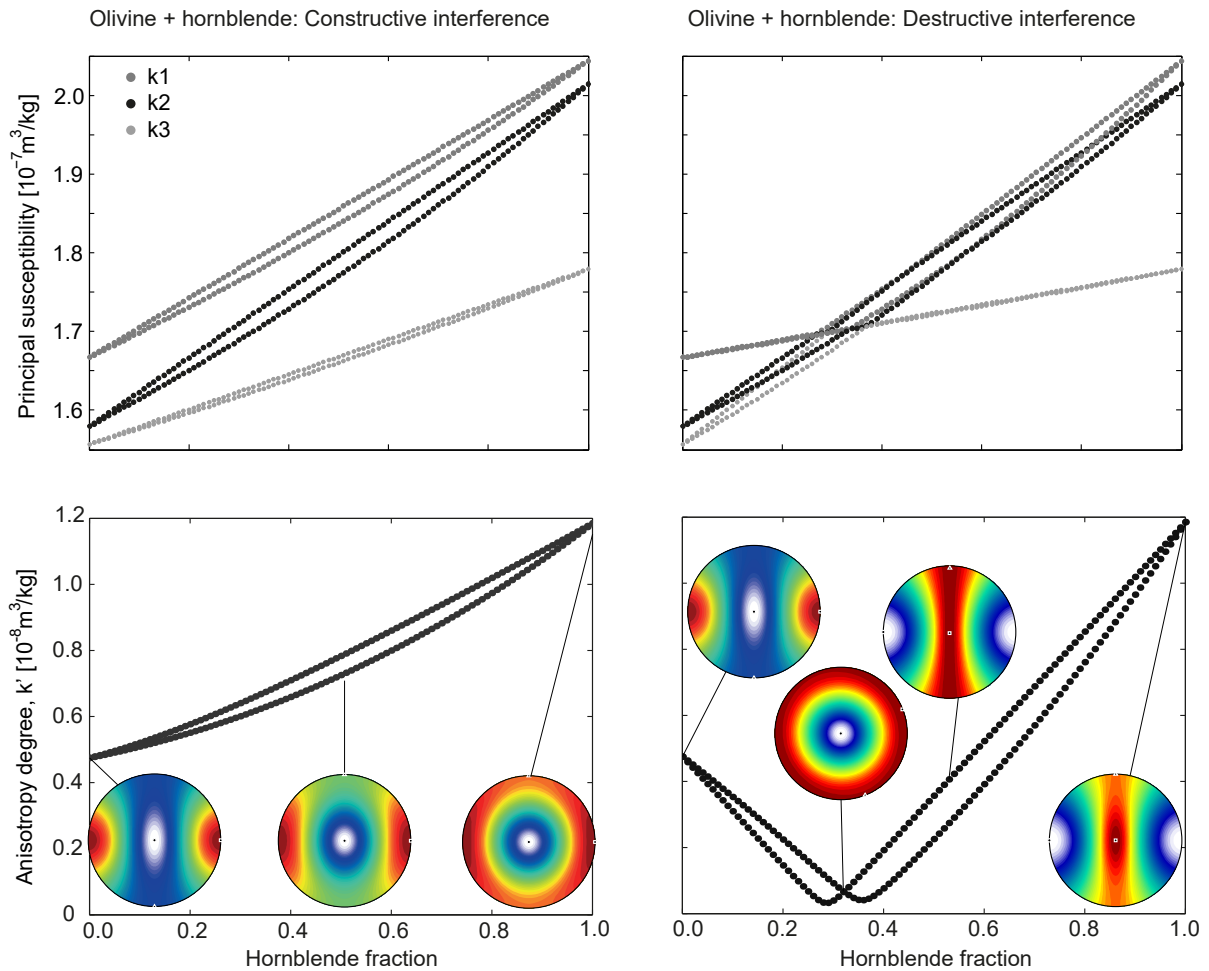


Fig. 3. Synthetic example for positive and negative interference between olivine and hornblende AMS. Modified after a conference presentation by Biedermann et al. (2014b): Upper and lower dotted lines show Voigt and Reuss averages, respectively, for the modelled principal axes and anisotropy degree (Reuss, 1929; Voigt, 1928).

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