

Inside...

Visiting Fellows' Reports

2

Current Articles

6

... and more throughout!

The IRM at 25: A Quarter Century of Community-Based Research and Education

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 Institute for Rock Magnetism

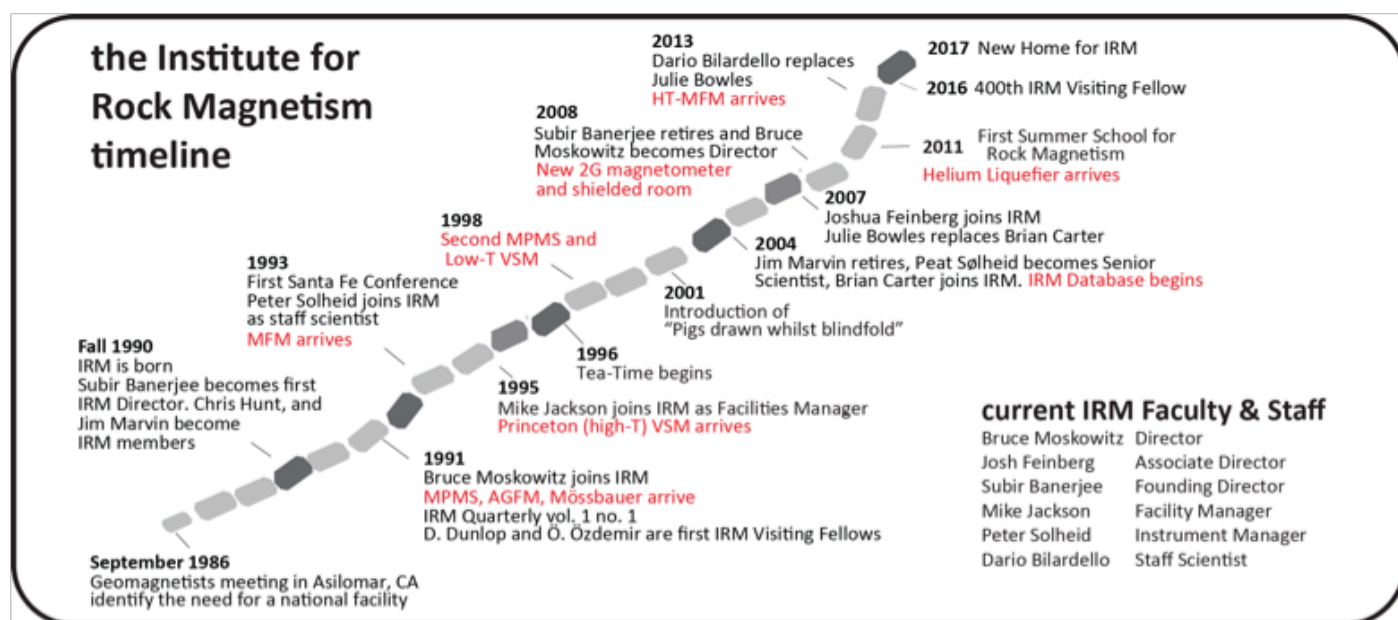


Fig 1. 30-year chronology of IRM development, schematically represented by a chain of bacterial magnetosomes.

IRM Origins and Development

In 1986, at an Asilomar meeting on the future of paleo- and rock magnetism organized chiefly by Subir Banerjee, 28 researchers began discussing an idea for a center where researchers in earth science and allied fields in the physical sciences could share ideas and have access to advanced instrumentation for magnetic material characterization and experimentation. Sensitive magnetic measurements would make it possible to determine the compositions, concentrations, particle sizes and orientation distributions of the trace magnetic phases in terrestrial and extraterrestrial materials. This in turn would improve our understanding of the origins and significance of the natural remanence carried by them, as well as providing valuable information on the physical and chemical processes involved in their formation and subsequent alteration.

The first attempt to fund such a center (originally called the National Center for Rock Magnetism) was in 1988 through a new initiative at NSF called the Science and

Technology Centers (STC) program. The proposal was reviewed favorably and it actually made it into the final round of competition, and when the review panel came for a site evaluation they were told by the governor of Minnesota that "we must have more rock magnetism." However the proposal was not funded in the end. Two Earth Science related STCs that did get funded in those early days of the program were by our colleagues in mineral physics (Center for High-Pressure Research at Stony Brook, NY) and seismology (Southern California Earthquake Center at USC, CA). A proposal for a more modest facility (now called the Institute for Rock Magnetism) proved to be much more successful in 1990, funded by the Instrumentation and Facilities program in the Division of Earth Sciences at NSF, the W. M. Keck Foundation, and the University of Minnesota. In hindsight it was probably lucky that we did not get the STC funding back in 1988 because the lifetime STC funding for individual centers was limited to 11 years total.

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Visiting Fellow Report

Anisotropy of magnetic susceptibility and remanence in layered intrusions – Rein-fjord Ultramafic Complex and Bjerkreim Sokndal Layered Intrusion

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Introduction

Magnetic fabrics are important proxies for rock textures (Borradaile and Jackson, 2010). Magnetic anisotropy also deflects magnetization and influences magnetic anomalies or paleomagnetic data (Clark, 1997; Fuller, 1963). We are interested in characterizing and understanding magnetic fabrics in two layered intrusions: the Rein-fjord Ultramafic Complex in the Seiland Igneous Province, Northern Norway, and the Bjerkreim Sokndal intrusion, Rogaland, Southern Norway.

The Rein-fjord Ultramafic Complex (RUC) consists of > 25 000 km³ ultramafic material, emplaced at 25-30 km depth (Roberts et al., 2006). It is currently not understood how this large amount of ultramafic melt could have been emplaced in the crust. Anisotropy of magnetic susceptibility (AMS) is used to better understand the rock fabric, and assist the geological interpretation. It is thus essential to separate paramagnetic and ferromagnetic fabrics, which are carried by olivine and pyroxene, and by magnetite (\pm sulfides), respectively.

The Bjerkreim Sokndal Layered Intrusion (BKSK) consists of plagioclase pyroxene cumulates, and the layering in the intrusion is associated with distinct magnetic anomalies. A strong negative anomaly, reducing the surface field by about half, is associated with one of the layers, MCU IVe' (McEnroe et al., 2001). This layer contains hemo-ilmenite and magnetite in addition to the paramagnetic minerals. The natural remanent magnetization (NRM) causing the negative anomaly is carried by hemo-ilmenite, which is expected to have a preferential orientation. We are interested in how the changing mineral fabric along the fold affects the observed anomaly at the surface.

Low-field AMS in both intrusions is strong ($P = k_1/k_3$ up to 1.45 in RUC, and up to 2.13 in BKSK) and the principal susceptibility directions appear to be related to geologic features. In the RUC, the maximum susceptibility direction is parallel to the main geological dip direction, and in the BKSK, the minimum susceptibility is normal to the trend of the layering, which corresponds to the trend of the negative anomaly. For a full interpretation of AMS data, it is however essential to know which minerals carry the AMS.

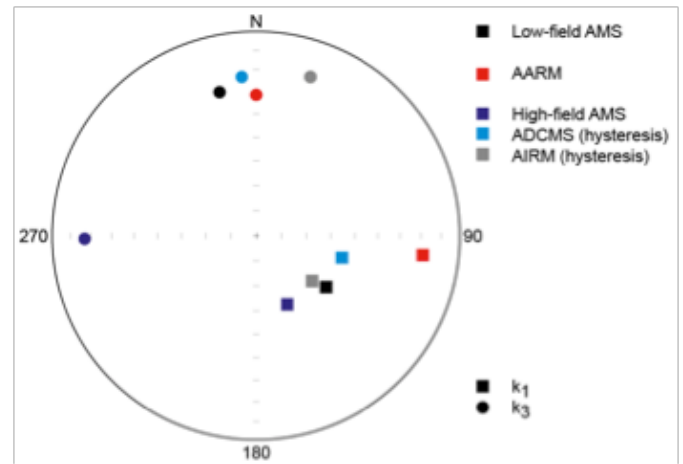


Figure 1. Maximum and minimum principal susceptibility and remanence directions of a representative sample from the RUC.

Methods

A series of experiments were conducted to isolate individual components of the magnetic fabric: High-field anisotropy measurements based on hysteresis loops, anisotropy of anhysteretic remanent magnetization (AARM), anisotropy of isothermal remanent magnetization (AIRM) and anisotropy of thermal remanent magnetization (ATRM). These measurements were complemented with rock magnetic experiments, i.e. acquisition of isothermal remanence up to 9 T, and field-cooled/zero-field-cooled remanence.

Results

High-field AMS and anisotropy of remanence have been measured on the Seiland samples (Fig. 1). The anisotropy of remanence and low-field AMS have similar orientations of maximum and minimum susceptibility axes, whereas the minimum axis of the high-field paramagnetic AMS is perpendicular to the minimum susceptibility of the low-field AMS.

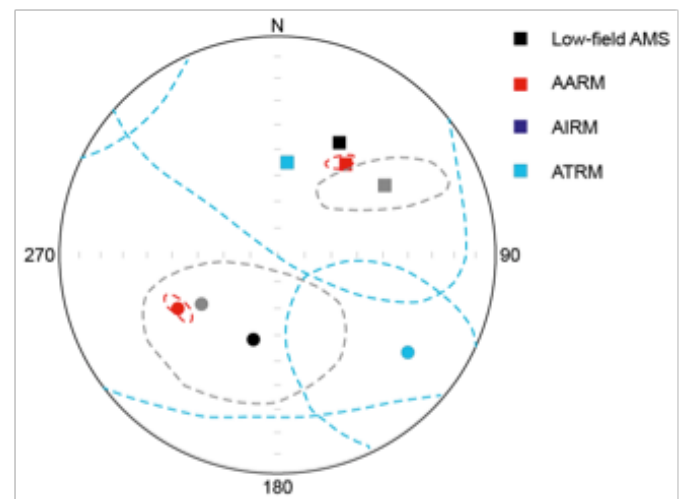


Figure 2. Maximum and minimum principal susceptibility and remanence directions of a representative sample from the Bjerkreim Sokndal layered intrusion. Dashed lines indicate confidence ellipses.

The samples from Bjerkreim-Sokndal were too

magnetic for high-field AMS measurements. AARM is strong and in most samples coaxial with low-field AMS (Fig. 2). AIRM and ATRM were hardly significant or not significant. This indicates that a low-coercivity mineral such as magnetite contributes a large part to the low-field AMS. The high-coercivity ilmenite was either not magnetized in the 1 T field used for AIRM, or does not show any anisotropy.

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Report from the 2015 Visiting Fellowship at the IRM

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Chalk rocks within the Dead-Sea-Fault plate boundary exhibit well defined magnetic fabrics. The rocks show both AMS and AARM fabrics (Figure 1), however those two differ in characteristics. The AARM fabric is held only by remanence and thus suggest preferred orientation of ferromagnetic grains.

In order to correctly interpret the magnetic fabrics, knowledge of the magnetic mineralogy is required. Therefore, the main goal for the IRM visit was to determine the magnetic mineralogy within the samples. For that, 15 samples were examined by: AF demagnetization curves, VSM hysteresis loops, thermal IRM demagnetization, high temperature susceptibility measurements and MPMS experiments.

The challenge was that the content of magnetic min-

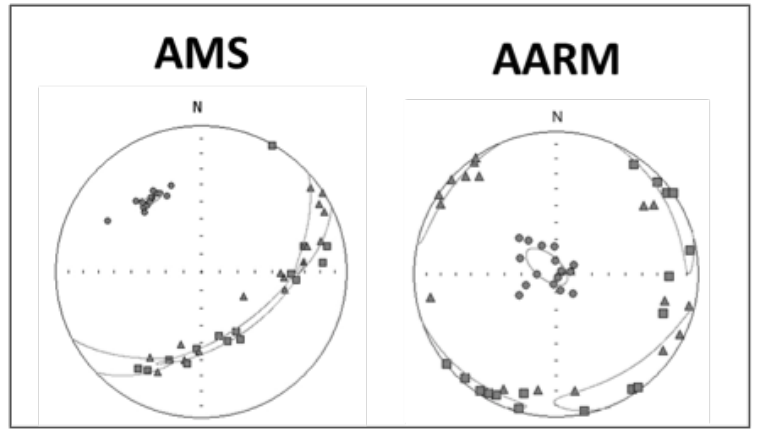


Figure 1. AMS and AARM fabrics of chalk within the Dead-Sea-Fault plate boundary.

erals is very low, and is masked by large contents of paramagnetic and diamagnetic minerals. For that, all the experiments and sample handling had to be very careful and precise. For example, to increase the accuracy, the VSM runs were set for high averaging time and 8 repetitions, and the sample holder was separately measured and then subtracted.

The main results are presented in Figure 2. The AF demagnetization curves indicates soft minerals, as most of the moment is lost between 10-35 mT (Figure 2a). The shape of the curve suggest sub-micron grain size (Figure 2a). The raw data of the VSM (Figure 2b) indicates the dominance of paramagnetic minerals as ferromagnetic component is negligible. The slope corrected curve (Figure 2c) however, reveals the existence of very low corecivity ferromagnetic phase. The shape of the loop is consistent with super-paramagnetic grains (Figure 2c).

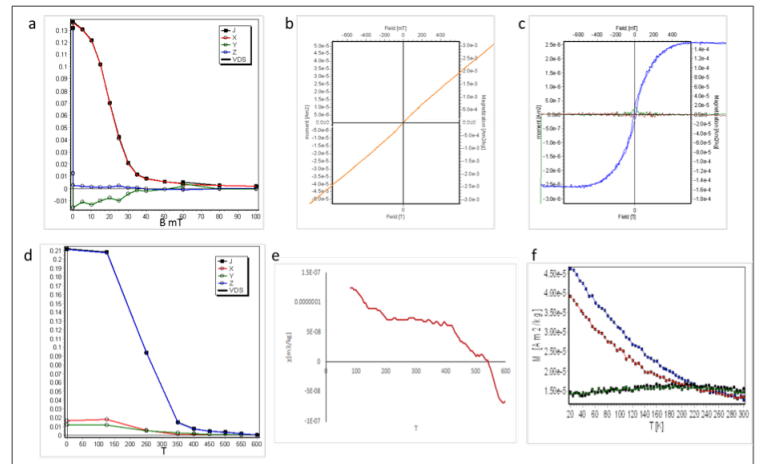


Figure 2. Main results of the experiment conducted at the IRM. a) AF demagnetization curves. b) Raw data of VSM loops. c) Slope corrected VSM loops. d) Thermal IRM demagnetization, the IRM was first acquired 1T along x, 300mT along y and then 100mT along z. e) High temperature susceptibility measurements. f) MPMS experiments.

Thermal demagnetization (Figure 2d) indicates that most of the initial magnetization is along z axis, also suggesting soft minerals (the IRM was first acquired 1T along x, 300 mT along y and then 100 mT along z). Most

of the moment loss is between 125°-350° as typical for sulfides. The remaining magnetization is lost after 500° as typical for magnetite-maghemite. High temperature susceptibility (Figure 2e) shows two drops during heating, at 100°-200° and 400°-600°. After ~550° the susceptibility in negative, indicates diamagnetic component.

The MPMS (Figure 2f) experiment shows no phase transitions at cooling, suggest the absence of magnetite or pyrrhotite. The curve indicates very little change at cooling, monotonic loss of moment on warming and reversibility during cooling warming cycle.

Examination of the whole results suggest that magnetic minerals are fine grained, probably super-paramagnetics, with very low corecivity. The most likelihood mineralogy emerges is a combination of greigite and oxidized magnetite.

Particle size dependent microbial oxidation and reduction of magnetite

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Iron (Fe) minerals are ubiquitous in the environment and commonly exist in combination with a number of different Fe-metabolizing bacteria including both Fe(II)-oxidizing and Fe(III)-reducing bacteria [1]. The formation of magnetite by magnetotactic bacteria [2] or dissimilatory iron(III)-reducing bacteria (e.g. *Geobacter sulfurreducens*) [3] has been known for several decades. However, the potential for bacteria to further interact with the magnetic mineral once formed is relatively unknown despite the fact that microbial-mineral interactions play a crucial role in the mobility of metals and nutrients, and for establishing certain redox conditions in the environment. We recently showed that Fe-metabolizing bacteria can use magnetite nanoparticles as an electron sink or source depending upon the redox conditions present, and likened the magnetite to a natural battery (biogeochemical battery) [4]. In that study, we incubated magnetite nanoparticles in the light with the phototrophic Fe(II)-oxidizing bacteria *Rhodospseudomonas palustris* TIE-1 [5] followed by incubation in the dark with Fe(III)-reducing bacteria *Geobacter sulfurreducens*. We used several techniques including magnetic susceptibility, x-ray diffraction (XRD) and Mössbauer spectroscopy to observe changes to the redox state of the mineral. This process of light/dark incubation was repeated, showing that the process of microbial oxidation and reduction of magnetite could be cycled. Many open questions emerged as a result of this research, notably the potential effect that particle size would have on the microbe-mineral interaction. Furthermore, we wanted to know if these changes, which were detectable through several mineralogical methods, could be observed using

low temperature magnetic measurements. This led to our application to carry out measurements at the Institute for Rock Magnetism (IRM).

Two different types of magnetite were synthesized, denoted nano-ctrl (d~12 nm) [6] and micro-ctrl (d~200-300 nm) [7] (Figure 1). These magnetite samples were incubated in the presence of either Fe(II)-oxidizing bacteria *Rhodospseudomonas palustris* TIE-1 (denoted nano-ox or micro-ox), or the Fe(III)-reducing bacteria *Geobacter sulfurreducens* (denoted nano-red or micro-red). After several days of incubation, the magnetite was dried for the magnetic measurements. Low temperature measurements were performed using a Quantum Design MPMS-2 superconducting quantum interface device (SQUID) magnetometer. Field cooled-zero field cooled saturation isothermal remanent magnetisation (FC-ZFC-SIRM) curves were obtained together with room temperature SIRM (RTSIRM) protocols. The data collected has been submitted for publication.

FC-ZFC-SIRM measurements of nano-ctrl showed a relatively steep decrease in magnetization during warming from 20 K to ~110 K followed by a more gradual decrease as temperature increased to 295 K. Micro-ctrl showed much more characteristic behaviour of bulk magnetite with a gradual decrease in magnetization on warming until above 100 K at which point the magnetization sharply decreased before plateauing above ~120 K. The rapid change in the slope of magnetization for both nano-magnetite and micro-magnetite samples can be attributed to the Verwey transition (T_v) [8]. T_v is commonly described to occur at ~120 K, and any changes in the transition temperature can be thought of as being due to oxidation (maghemitization) of magnetite. It was seen in the data that T_v decreased by ~5 K for nano-ox in comparison to nano-ctrl, confirming that the sample was partially oxidized by the bacteria. No such decrease in T_v could be observed for micro-ox compared to micro-ctrl which suggests oxidation did not occur. Changes in the RTSIRM curves were most apparent for the nano-magnetite series of samples. Whilst the cooling curves for nano-ctrl, nano-ox and nano-red share similar characteristics, the warming curve for nano-ox shows a much more dramatic drop in magnetization above ~220 K and continue to fall to 80% of the initial magnetisation before cooling and is thought to be an effect of microbial Fe(II) oxidation. These results confirm that nano-magnetite was more susceptible to oxidation than micro-magnetite. This behaviour was expected due to the increased surface area to volume ratio of nano-magnetite enabling a larger portion of the Fe to be available to the Fe(II)-oxidizing bacteria. The measurements were, however, unable to confirm such a particle size based effect for the Fe(III)-reducing bacteria as the nano-red and micro-red samples showed similar characteristics to the nano-ctrl and micro-ctrl samples respectively.

In summary, these magnetic-based approaches have helped to further develop an understanding of how Fe-metabolizing bacteria can utilize magnetite for growth and respiration. This has several important consequences in terms of using magnetic measurements to search for

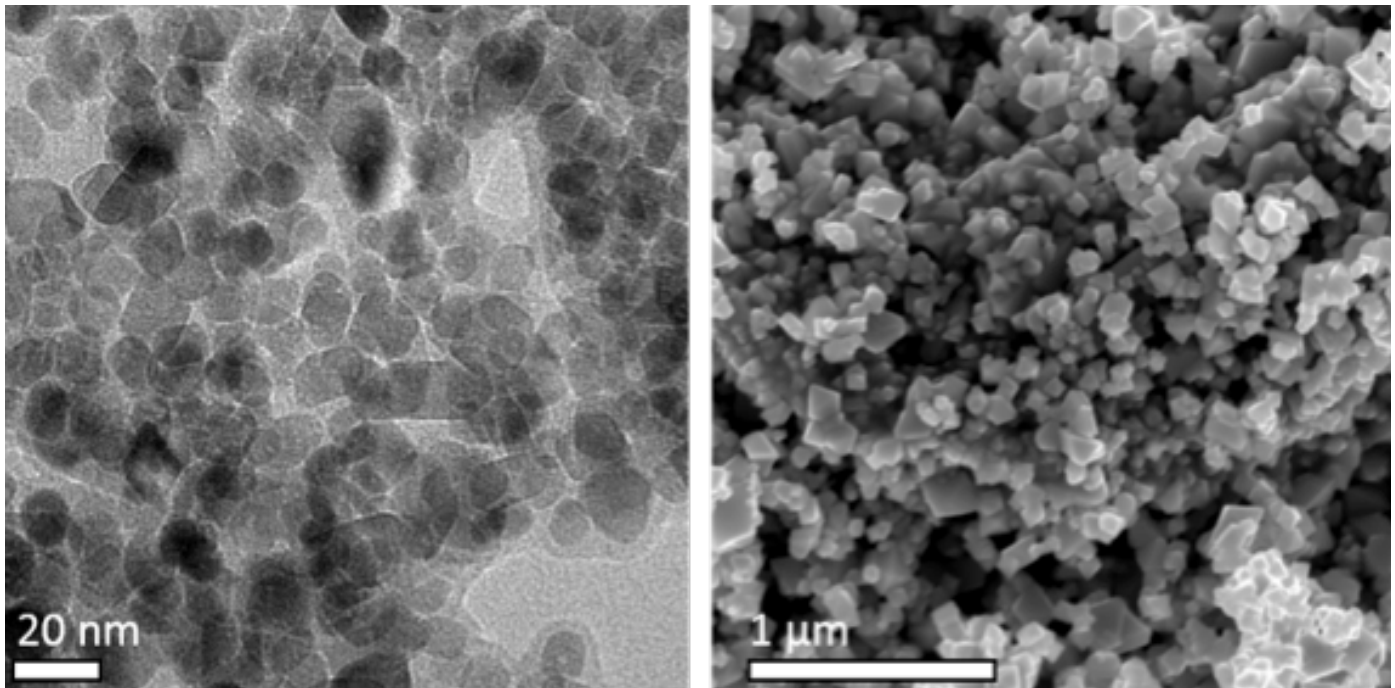


Figure 1. Electron micrographs of (a) nano-magnetite and (b) micro-magnetite.

signatures of current or past microbial activity in soils, sediments and potentially even the rock record. For example, increases or decreases in magnetization of environmental samples could suggest microbial Fe(III) reduction or Fe(II) oxidation respectively rather than simply being due to increases or decreases in the overall magnetic content of a sample. These type of measurements could also potentially be used as evidence of previous changes in redox conditions which favour either type of Fe-metabolizing bacteria.

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The next **Visiting Fellow Application Deadline** is April 30th! Find details at <https://irm.umn.edu>

Current Articles

A list of current research articles dealing with various topics in the physics and chemistry of magnetism is a regular feature of the IRM Quarterly. Articles published in familiar geology and geophysics journals are included; special emphasis is given to current articles from physics, chemistry, and materials-science journals. Most are taken from ISI Web of Knowledge, after which they are subjected to Procrustean culling for this newsletter. An extensive reference list of articles (primarily about rock magnetism, the physics and chemistry of magnetism, and some paleomagnetism) is continually updated at the IRM. This list, with more than 10,000 references, is available free of charge. Your contributions both to the list and to the Current Articles section of the IRM Quarterly are always welcome.

Archeomagnetism

- Donadini, F., V. Serneels, L. Kapper, and A. El Kateb (2015), Directional changes of the geomagnetic field in West Africa: Insights from the metallurgical site of Korsimoro, *Earth and Planetary Science Letters*, 430, 349-355.
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Subir (left) and Bruce (right) ca 1989.

The idea hatched four years earlier at Asilomar became reality in 1990, when the Institute for Rock Magnetism (IRM) was established as a shared resource for the GPE and broader research communities, providing instruments to study the magnetism of rocks, sediments, biological materials and synthetic analogs. This is accomplished with a suite of instruments that measure magnetization from 2-1000 K, in DC fields up to 5 T and AC fields up to 10 kHz, and a low-temperature probe (20-300 K) for vector remanence measurements. These magnetometric capabilities are complemented by Mössbauer spectrometers (4.2-300K, 0-6.5T) and a high-temperature magnetic force microscope (Tmax~ 673 K). A unique aspect of the IRM was that it allowed for routine measurements below 300 K and provided new ways of "seeing" magnetism. This has enabled researchers to study magnetic behavior through magnetic ordering temperatures, crystal phase transitions, and blocking temperatures, providing new insights into mineral magnetism as well as developing new methods to interpret the magnetism of natural materials.

IRM Oversight and Planning

From the beginning the IRM has operated under the joint guidance of its local leadership (Fig 1) and an external Review and Advisory Committee (RAC), an international panel of leading researchers in paleomagnetism, rock magnetism and related fields (Table 1). To date, 38 of our esteemed colleagues have contributed their time, efforts and strategic vision to the RAC, helping to keep the IRM's instrumental capabilities aligned with the needs of the research community, and evaluating all Visiting Fellowship applications. Members currently serve 4-year terms, during which they typically review about 100 Visiting Fellow proposals, and participate in two full-day meetings at the IRM as well as additional shorter sessions held in conjunction with Santa Fe and AGU conferences. The active participation of these distinguished colleagues has been an essential ingredient in maintaining the IRM as an effective resource for the GP research community.

Visiting Fellows

The main access to the IRM by the research community is through the Visiting Fellowship (VF) program. To date, 388 Fellowships have been awarded to students, post-docs and senior researchers representing 157 institutions in the US and 30 other countries. Nearly half of these Fellowships have been awarded to researchers from universities with strong programs in paleomagnetism and rock magnetism (Toronto-Erindale 26; Michigan 14; UC-Santa Cruz 13; LMU München 12; Wisconsin-Madison 9; IPGP 8; Lehigh 8; ODP 8; NGU Trondheim 7; Michigan Tech 7; SIU Carbondale 7; Massachusetts-Amherst 7; New Mexico 7; Rochester 7; ETH Zürich 6; Imperial College London 6; UC-Berkeley 6; UNAM 6; Florida 6; Oklahoma 6; Yale University 6). Nearly 50% of Visiting Fellows have been students.

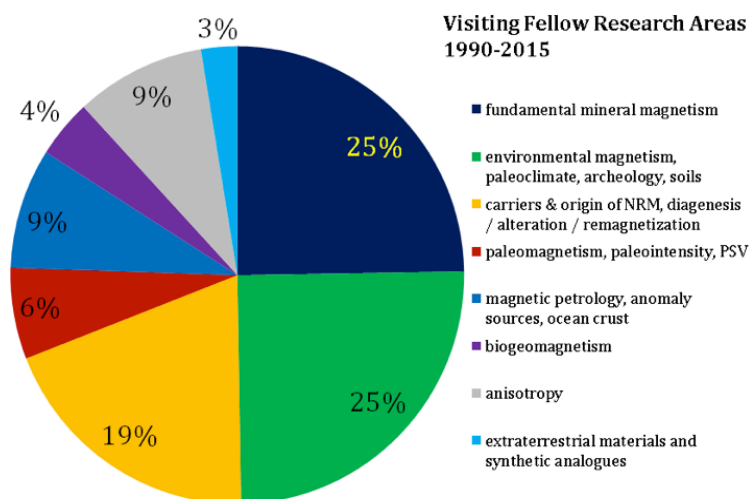


Fig 2. General breakdown of research areas investigated by Visiting Fellows at the IRM over the past 25 years.

Visiting Fellow research projects have covered a remarkably broad range of scientific fields, problems, geologic settings and materials (Fig. 2). Pseudotachylites, horse-spleen ferritin, the Chelyabinsk meteorite, Armenian obsidian, Chinese loess, Pittsburgh tree leaves, Brazilian termite nests, Middle Eastern archeological copper slags, Martian meteorites, the Vredefort impact structure, Antarctic shelf sediments, Pangean red beds, cultured and wild magnetotactic bacteria, oil sands, Paleocene-Eocene boundary sediments, blood diamonds, “snowball Earth” sediments and cap carbonates... well, you get the idea. The variety of scientific problems addressed is just as great as the variety of materials, including the origins and evolution of the magnetic fields of the terrestrial planets; tectonic plate motions and mantle dynamics; changing environmental conditions over the Earth’s surface through time; and the microscale physical and chemical processes by which these large-scale phenomena leave their imprint in the natural remanence of geomaterials and in their magnetic mineral characteristics. Fundamental studies of mineral magnetism (magnetic ordering and intrinsic properties, intraparticle spin structures, etc) and rock magnetism (magnetic behavior

of interacting and noninteracting populations with distributions of particle size and of easy-axis orientation) represent the largest single research area for Visiting Fellows, followed closely by environmental studies and characterization of paleomagnetic remanence carriers (Fig 2). The total output of visiting and in-house researchers amounts to about 800 peer-reviewed publications to date.

IRM Conferences, Workshops, and Summer School

The biennial Santa Fe meetings feature a limited number of presentations and extended informal discussions on the current state and future trends in magnetic research. The series began at the Santa Fe campus of St John’s College in 1992, and has returned there nine times subsequently; the meeting has twice been transplanted to other venues (Erice, Sicily in 2002 and Cargèse, Corsica in 2008) to enable greater participation by our European colleagues and to benefit from their particular areas of strength in research. As described in the previous issue of the Quarterly, we plan to shift the future meeting schedule to odd-numbered years, and current plans are to hold the next meeting in Australia in 2017.



St John’s College in Santa Fe, site of ten conferences on rock magnetism and its applications.

The biennial Summer Schools in Rock Magnetism offer an intense 10-day instruction in rock magnetism theory and hands-on lab training to graduate students and post-doctoral researchers in the geosciences. Three Summer Schools have been held to date (2011, 2013 and 2015) and 64 students from 44 different academic institutions in 18 countries have completed the course. The fourth is scheduled for June 1-10, 2016.

IRM Future Developments

As described in the previous issue of the Quarterly,

a big change is in store for our facility: after 25 years of operation in the Shepherd Labs building, we will be relocating to a completely renovated space in the new home of the Earth Science Department and parts of the School of Physics and Astronomy at the University of Minnesota. We anticipate that this relocation will significantly improve facility operations through increased efficiency of spatial arrangements, as well as by integrating IRM more closely with the other geoscience research groups on campus, providing greater access by facility users to the intellectual and laboratory resources of the entire department. In addition, our lab will be on the second floor and our neighbors will include theoretical physicists and the Department of Astrophysics.

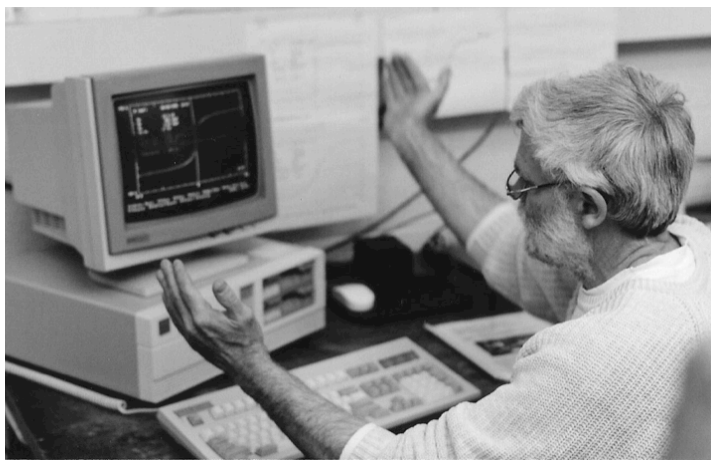
All Past and Present IRM RAC members:

Stefanie Brachfeld (Montclair State U.) 2012-present
 Laurie Brown (UMass Amherst) 2015-present
 Bob Butler (U. of Arizona) 1990-1996, chair 1994-1996
 Jim Channell (U. of Florida) 1998-2004
 Brad Clement (IODP) 2013-present, chair 2015-present
 Rob Coe (UC-Santa Cruz) 2001-2006, chair 2002-2006
 Cathy Constable (Scripps) 2009-present, chair 2010-2013
 Dan Dahlberg (U. of Minnesota, Physics) 1992-2000
 David Dunlop (U. of Toronto) 1990-1996, chair 1990-1994
 Mike Fuller (U.C. Santa Barbara) 1990-1994
 Jeff Gee (Scripps) 2005-2010
 John Geissman (U. of New Mexico) 2006-2009
 Sue Halgedahl (U. of Utah) 1994-2000
 Richard Harrison (Univ of Cambridge) 2015-present
 Friedrich Heller (ETH-Zurich) 1996-2002
 Dennis Kent (Lamont-Doherty) 1994-1998
 John King (U. of Rhode Island) 1994-2000, chair 1996-2000
 Ken Kodama (Lehigh U.) 1996-2002, chair 2000-2002
 France Lagroix (Inst. de Physique du Globe de Paris) 2015-present
 Cor Langereis, (U. of Utrecht) 2002-2006
 Chris Leighton (U. of Minnesota, Chem. Eng. & Mat. Sci.) 2008-2014
 Chad McCabe (Louisiana State U.) 1992-1994
 Suzanne McEnroe (Geol. Survey of Norway) 2010-2014, chair 2013-2014
 Ron Merrill (U. of Washington) 2001-2004
 Özden Özdemir (U. of Toronto) 2005-2010
 Yongxin Pan (Chinese Academy of Sciences) 2011-2014
 R. Lee Penn (U. of Minnesota, Chemistry) 2014-present
 Rich Reynolds (USGS) 1990-1994; 2010-2012
 Andrew Roberts (Southampton; Australian National U.) 2009-2013
 Pierre Rochette (CEREGE) 2002-2008
 Roy Roshko (U. of Manitoba, Physics) 2001-2008
 Mike Sharrock (3M Corp) 1990-1992
 Aleksey Smirnov (Michigan Tech.) 2013-present
 John Tarduno (U. of Rochester) 2005-2010, chair 2006-2010
 Lisa Tauxe (Scripps) 1998-2004
 Rob Van der Voo (U. of Michigan) 1990-1992
 Ben Weiss (MIT) 2010-present
 Toshi Yamazaki (Geol. Survey of Japan) 2007-2011

An IRM Photo Album



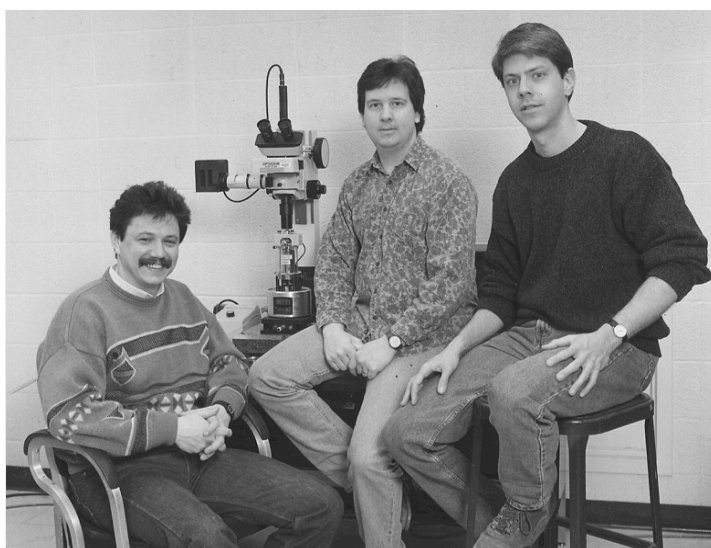
Pre-IRM post-doc Horst-Ulrich Worm operating VSM#1, ca 1989.



Jim Marvin measuring a rather sizeable hysteresis loop on the AGM, ca 1992.



IRM people in front of Shepherd Labs, 1993: (left to right) Subir Banerjee, Bruce Moskowitz, Stefanie Brachfeld, Chris Hunt, Peat Sølheid, Weiwei Sun, Sherry Foss, Gin Kletetschka, Li Xu, Jim Marvin, Scott Rubin.



IRM post-docs, 1994: Taras Pokhil, Bernie Housen, Rick Oches, with the first MFM.



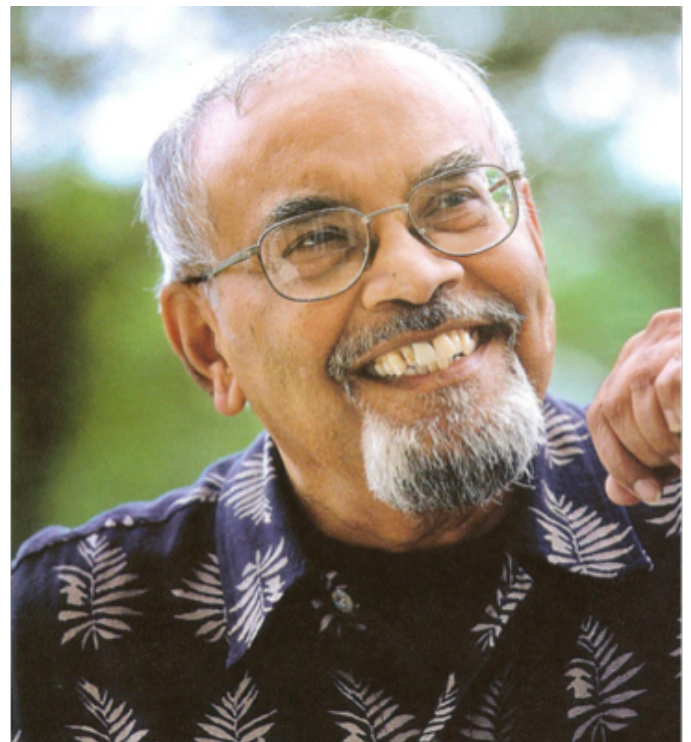
IRM staff/faculty 1994 (left to right) Chris Hunt, Jim Marvin, Bruce Moskowitz, Subir Banerjee.



A few of the 1996-1997 Visiting Fellows (clockwise from upper left): Andrei Kosterov (Univ. Montpellier), Adry Van Velzen (Oxford Univ.), Harald Peterman (Univ. Bremen), Luca Lanci (ETH-Zurich). Photos by Bernie Housen.



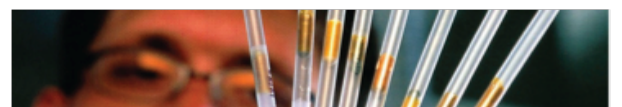
Panoramic self-portrait (gouache on digital scanner) by Christoph Eberhard Geiss (IRM Visiting Fellow 1993; IRM PhD student 1994-1999; IRM post-doc 1999-2001; IRM Visiting Fellow 2004, 2006, 2008).



Subir 2001.



David Williamson manning the AGM ca 1997.



Visiting Fellow Alex Zwing (LMU Munich) 2001.



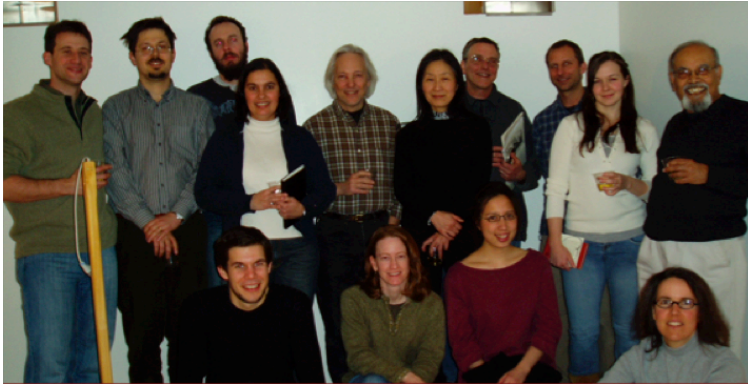
“Santa Erice” conference participants, 2002.



IRM resident researchers ca 2008. front: Jasmine Erbs, Thelma Berquó, Yifan Hu, Josh Feinberg, Mike Jackson. back: Amy Chen, Ioan (“Nono”) Lascu, Peat Sølheid, Subir Banerjee, R. Lee Penn, Jess Till, Julie Bowles.



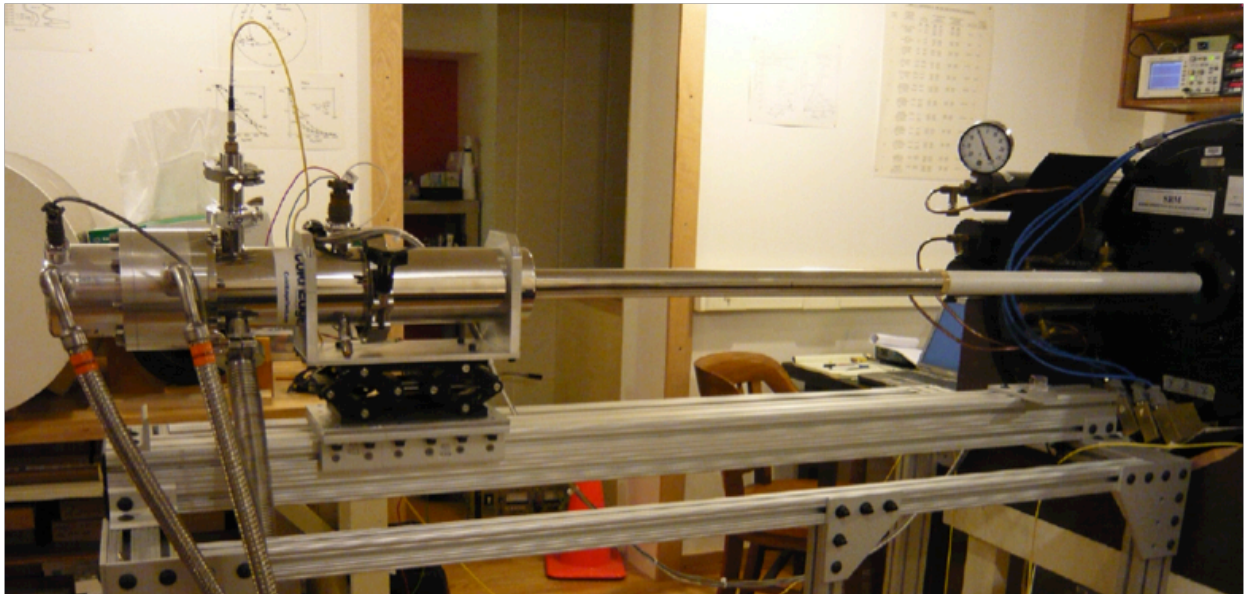
“Santa Cargèse” conference, 2008:
(left) outdoor poster session;
(right) the unforgettable boat trip to the Scandola Nature Preserve.



The newly completed shielded room 2008. front: Unidentified, Julie Bowles, Amy Chen, R. Lee Penn; (back): Josh Feinberg, Ramon Egli, Ioan (“Nono”) Lascu, Thelma Berquó, Bruce Moskowitz, Emi Ito, Mike Jackson, Peat Sølheid, Jess Till, Subir Banerjee.



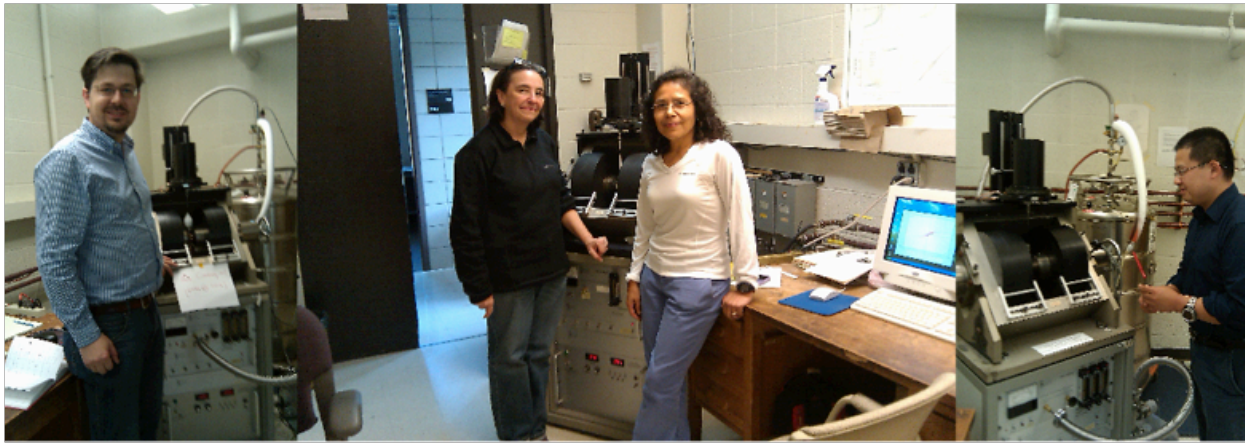
At Subir’s retirement fest 2008 (left to right) Yifan Hu, Brian Carter(-Stiglitz), Rachel (Carter-)Stiglitz, Taras Pokhil, Julie Bowles, Peat Sølheid, Max Brown, Jim Thill.



Low-temperature sample probe and 2G magnetometer, 2009.



8th Santa Fe Conference, 2010, field trip stop at outcrop of Bandelier Tuff.



The PMC VSMs, 2011: Ramon Egli (left) ; Ana Maria Soler and Beatriz Ortega (center); Huapei Wang (right).



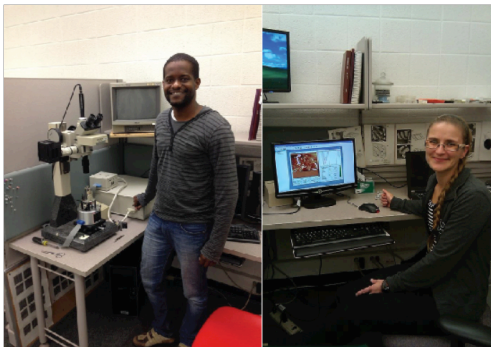
Prof Ed Nater, UM Dept of Soils, Water & Climate, leading students of the 1st Summer School in sampling a soil profile for magnetic analysis (2011).



Students of the 1st Summer School in a glacial pothole on a field trip to Interstate (MN-WI) Park (2011).



Field trip of the Second Summer School, 2013, Taylors Falls, MN: (left to right) Dario Bilardello, Sara Satolli, Peat Sølheid.



MFMI (Digital Instruments) operated by Edivaldo dos Santos (left, 2013) and Johanna Salminen (right, 2011).



At the 2nd Summer School, students Sara Satolli, Irina Seliverstov, Antonio Turtù and Samer Hariri were mesmerized by Nick Swanson-Hysell and the Kappabridge (2013).



Sketching pigs whilst blindfold at the closing ceremony of the Second Summer School, 2013.

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The IRM Quarterly

The *Institute for Rock Magnetism* is dedicated to providing state-of-the-art facilities and technical expertise free of charge to any interested researcher who applies and is accepted as a Visiting Fellow. Short proposals are accepted semi-annually in spring and fall for work to be done in a 10-day period during the following half year. Shorter, less formal visits are arranged on an individual basis through the Facilities Manager.

The *IRM* staff consists of **Subir Banerjee**, Professor/Founding Director; **Bruce Moskowitz**, Professor/Director; **Joshua Feinberg**, Assistant Professor/Associate Director; **Mike Jackson**, **Peat Solheid** and **Dario Bilardello**, Staff Scientists.

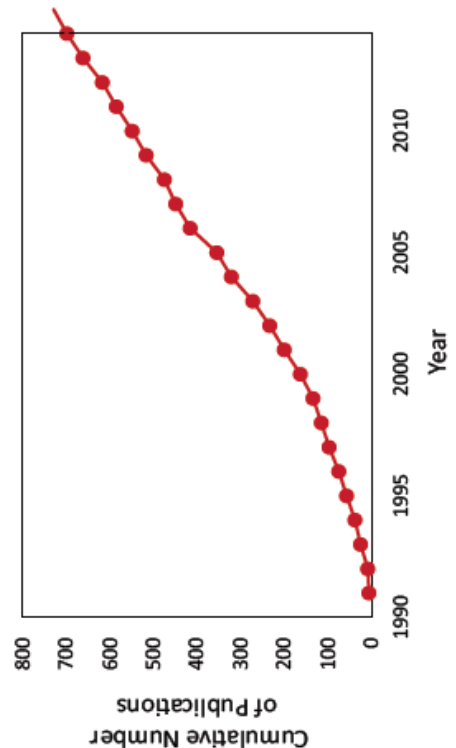
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The *IRM Quarterly* is published four times a year by the staff of the *IRM*. If you or someone you know would like to be on our mailing list, if you have something you would like to contribute (e.g., titles plus abstracts of papers in press), or if you have any suggestions to improve the newsletter, please notify the editor:

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